

Facilitating atmosphere oxidation through mantle convection

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Illustration by Richard Bizley

Acknowledgments



T. Gu
(Yale, now GIA)



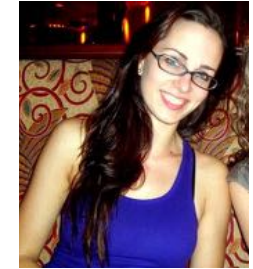
N. Creasy
(Yale)



M. Li
(ASU)



C. McCammon
(BGI)



J. Girard
(Yale)

Special thanks to:

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Facilitation of collaboration: CIDER and BurnMan

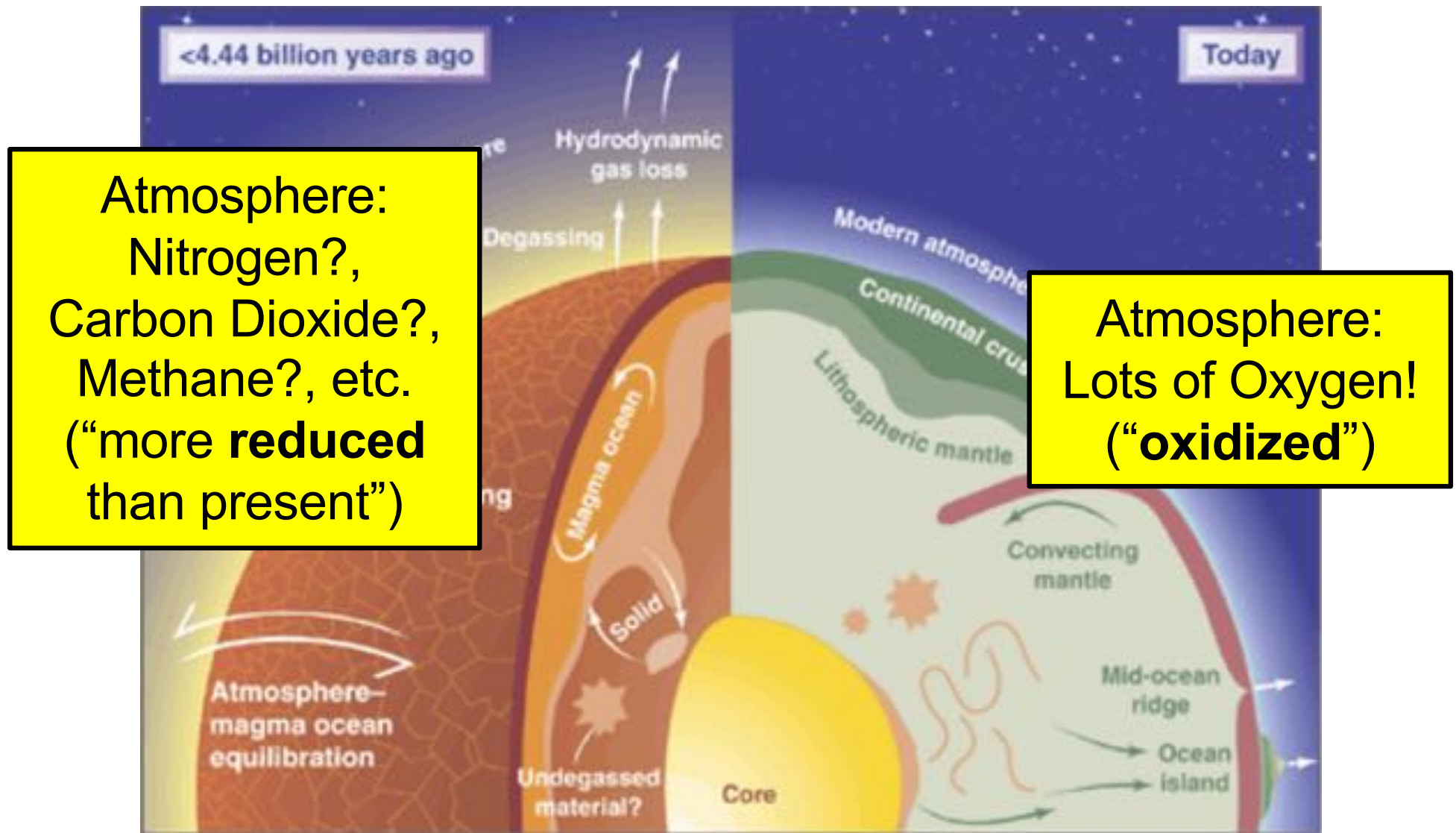
Outreach Travel: COMPRES

Funding provided by:

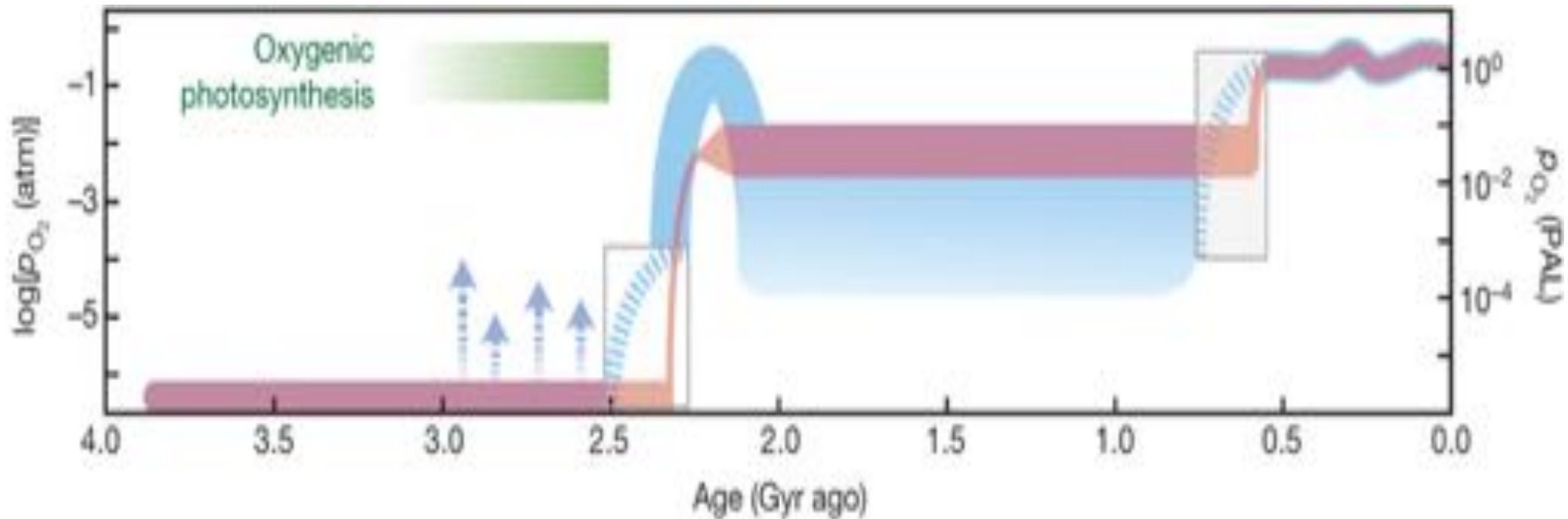
NSF CAREER (EAR-0955824), NSF GRFP



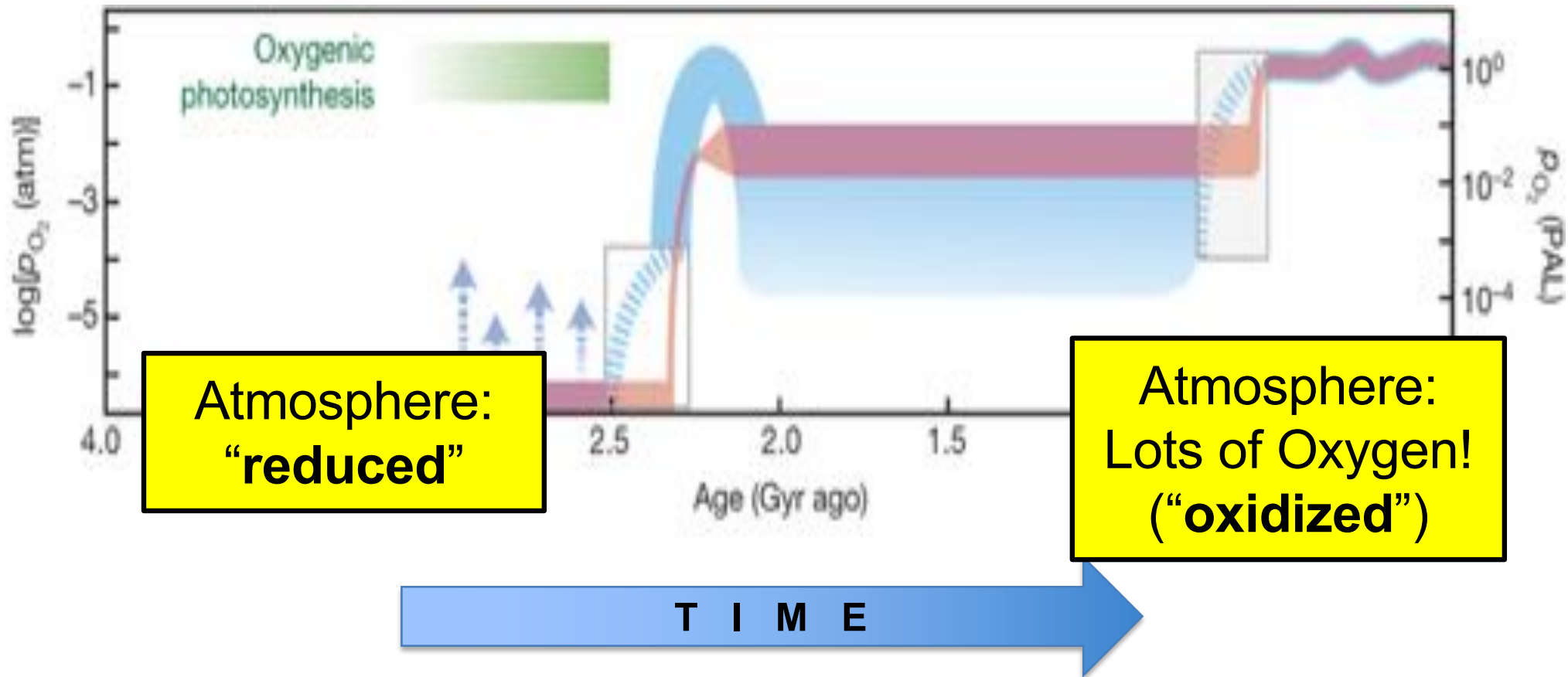
Evolution of O₂ in Atmosphere



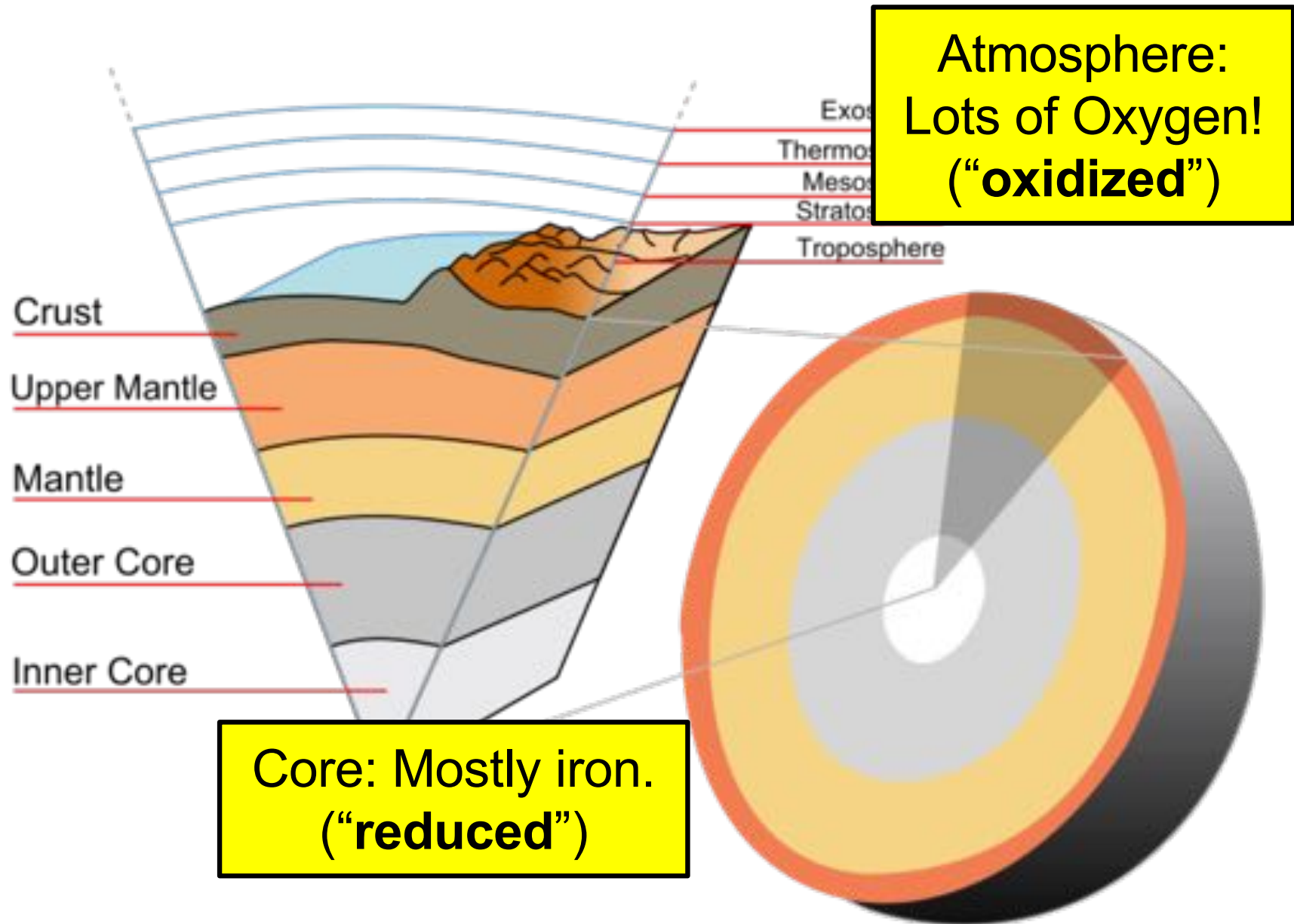
Evolution of O₂ in Atmosphere



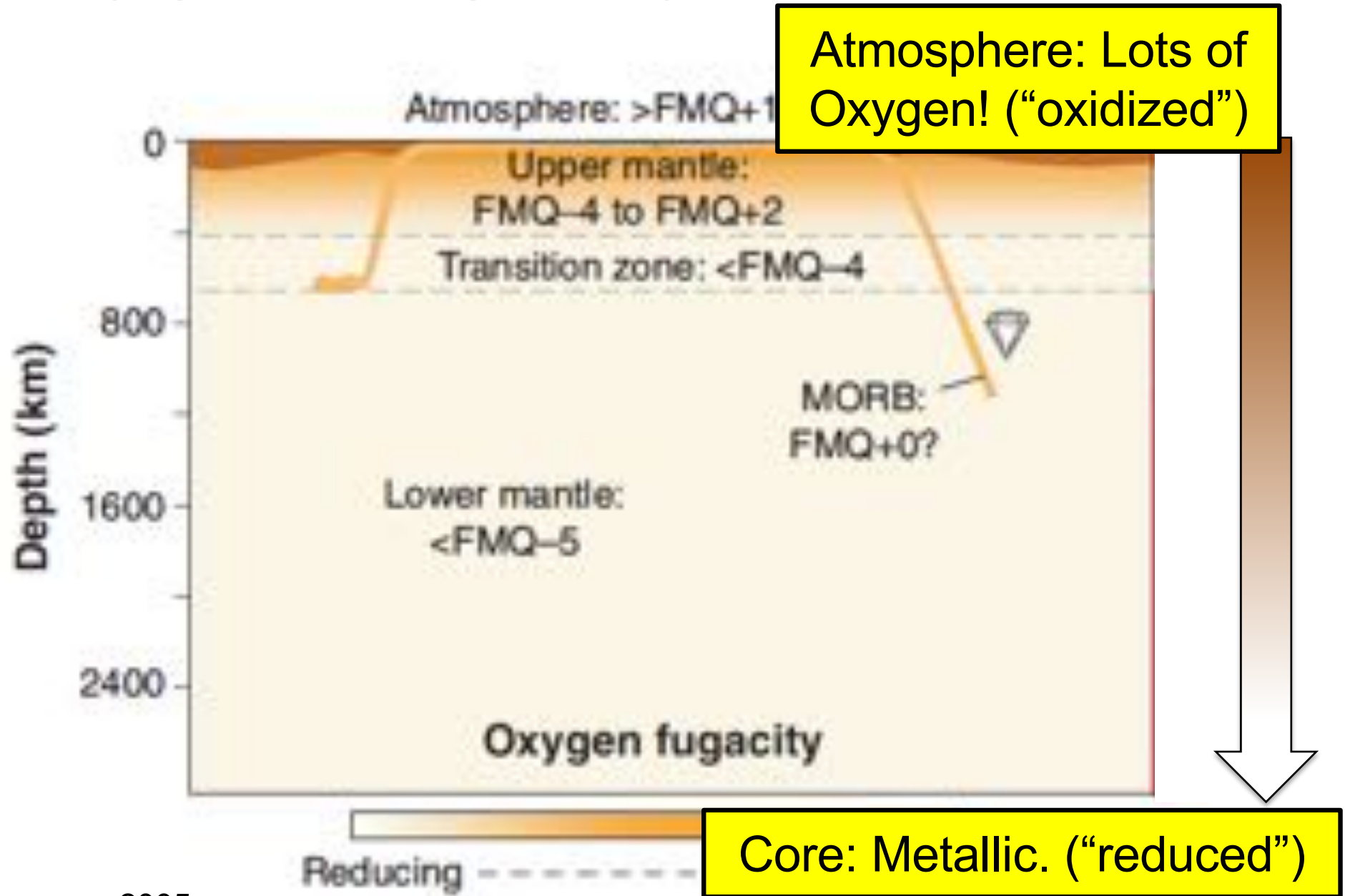
Evolution of O₂ in Atmosphere



Earth Structure



Oxygen Fugacity in the Mantle



Oxygen Fugacity in the Mantle?

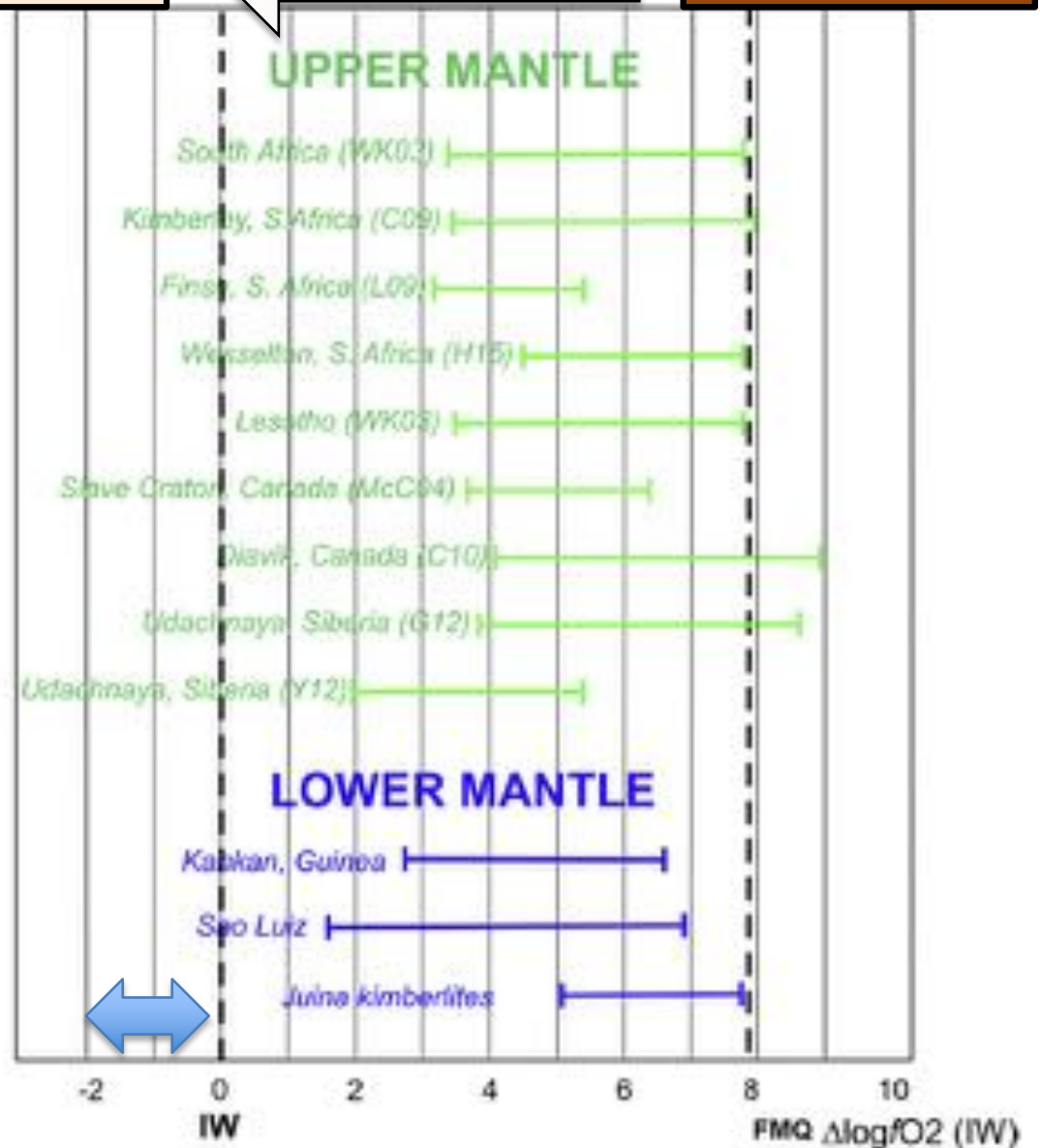
More reduced

More oxidized

What are diamond inclusions telling us about mantle redox spatially as well as temporally?

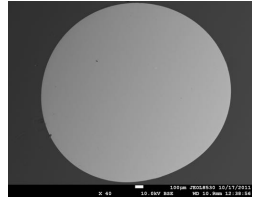


“Deep” LM inclusions



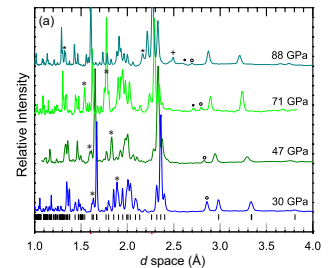
Methods

- **Synthesize** LM glasses: differ only in Fe^{3+} content

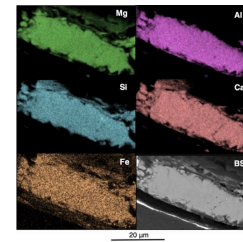


- **LHDAC** 

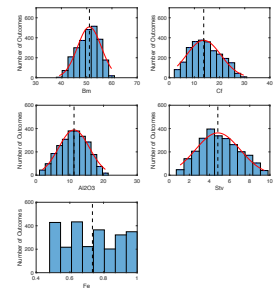
- **XRD**: identify minerals, equations of state



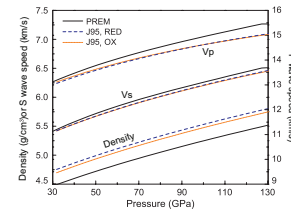
- **FIB** and **EPMA**: reveal compositions



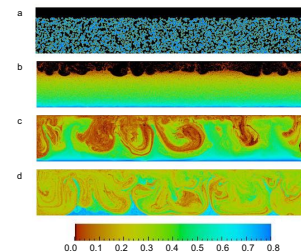
- **Monte Carlo Method**: determine abundance



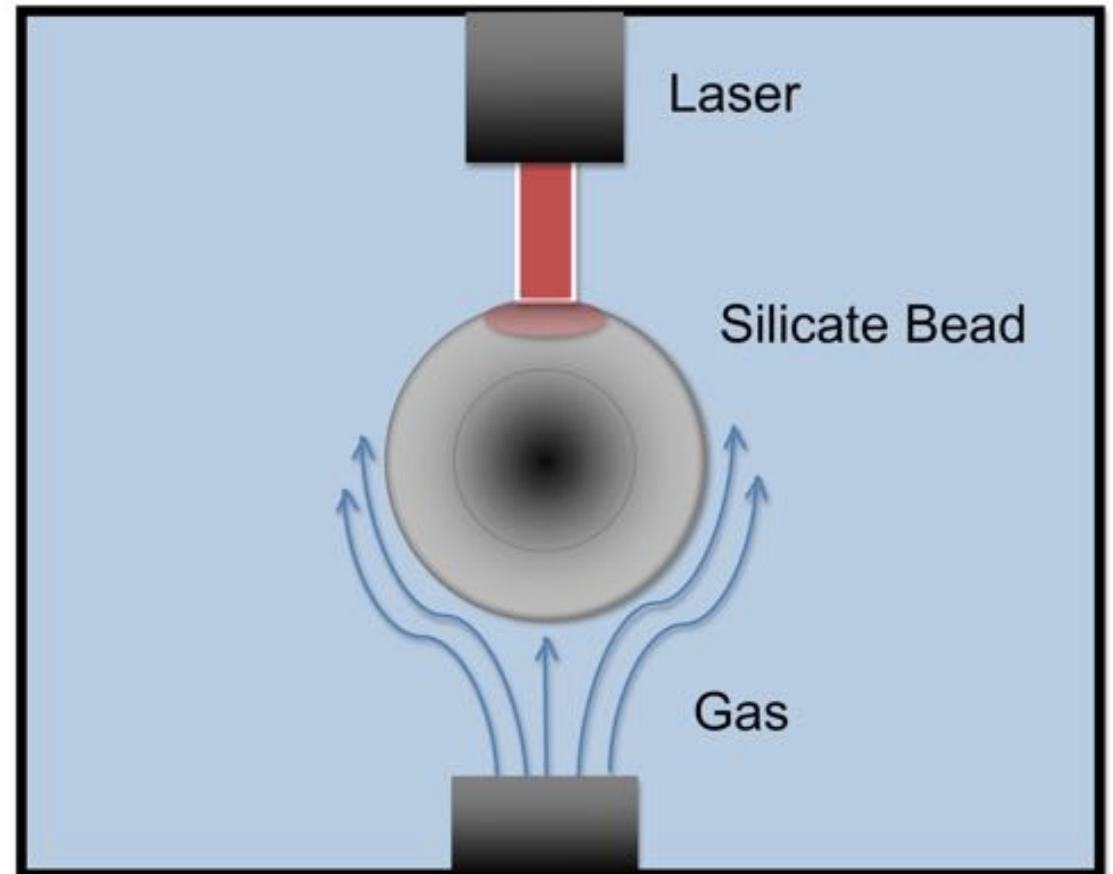
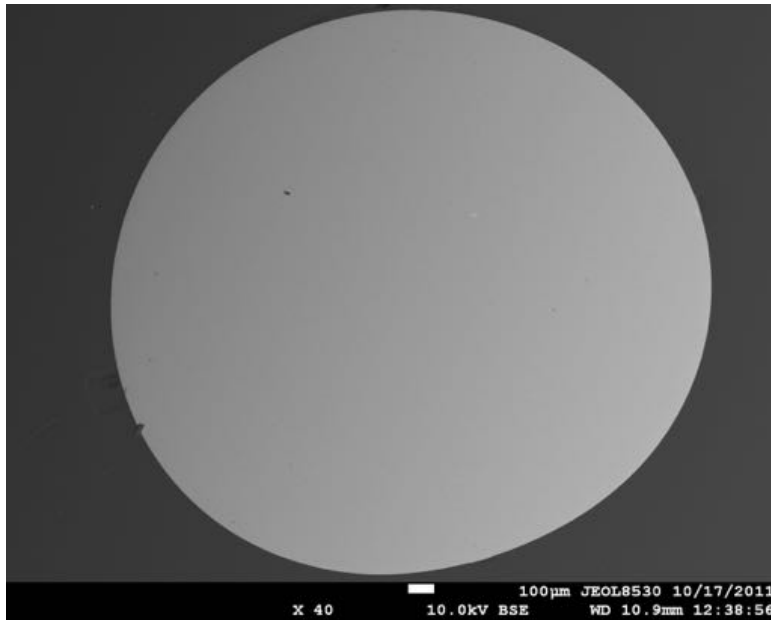
- **BURNMAN**: Compare with **PREM**



- **CITCOM** convection simulations



Starting Sample Synthesis



Temperature 2100-2600 K

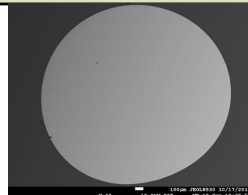
Levitation gas OXIDIZED: 100% O₂
REDUCED: 95% Ar, 4% CO, 1% CO₂

Composition **J95**: Javoy, 1995 (Enstatite Lower Mantle Model)
MIX1G: Hirschmann et al., 2003 (pyroxenite mixture and plausible OIB source)

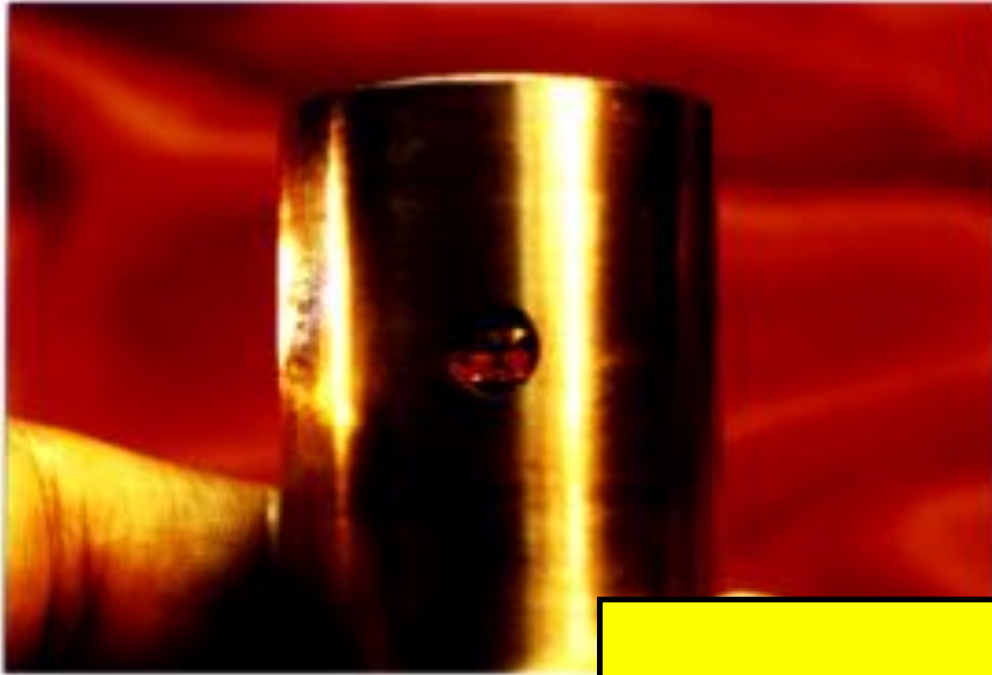
Mantle Compositions (in mol%)

MgO	SiO ₂	Al ₂ O ₃	CaO	FeO [^]	Molar Mass (g/mol)	Fe ³⁺ /Fe	Reference
36.9	47.3	1.2	1.7	12.9	54.79	0.30 (±0.04)	This study, J95 REDUCED [Gu et al., 2016]
37.1	47.3	1.2	1.6	12.7	54.70	0.35 (±0.04)	This study, J95 OXIDIZED [Gu et al., 2016]
37.1	47.7	1.1	1.7	12.5	54.67	NA	Enstatite Chondrite LM, [Javoy, 1995]
27.1	41.6	13.0	10.3	8.0	60.72	0.11 (±0.03)	This study, MIX REDUCED [Creasy et al., <i>in review</i>]
28.5	40.3	12.9	10.5	7.8	60.35	0.55 (±0.04)	This study, MIX OXIDIZED [Creasy et al., <i>in review</i>]
25.3	46.4	9.1	12.5	6.7	59.18	NA	MIX1G pyroxenite [Hirschmann et al., 2003]
49.4	39.2	2.3	3.3	5.8	51.83	NA	Pyrolite, [McDonough & Sun, 1995]

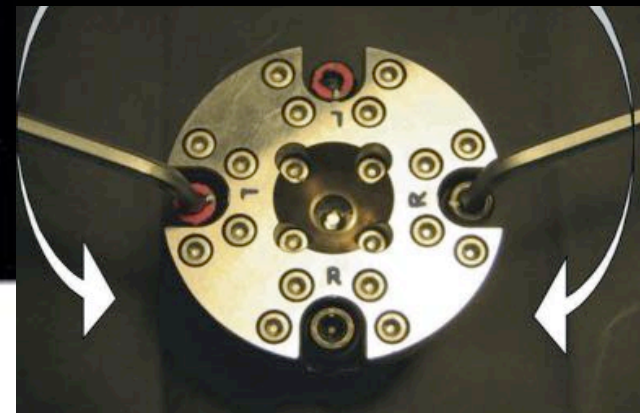
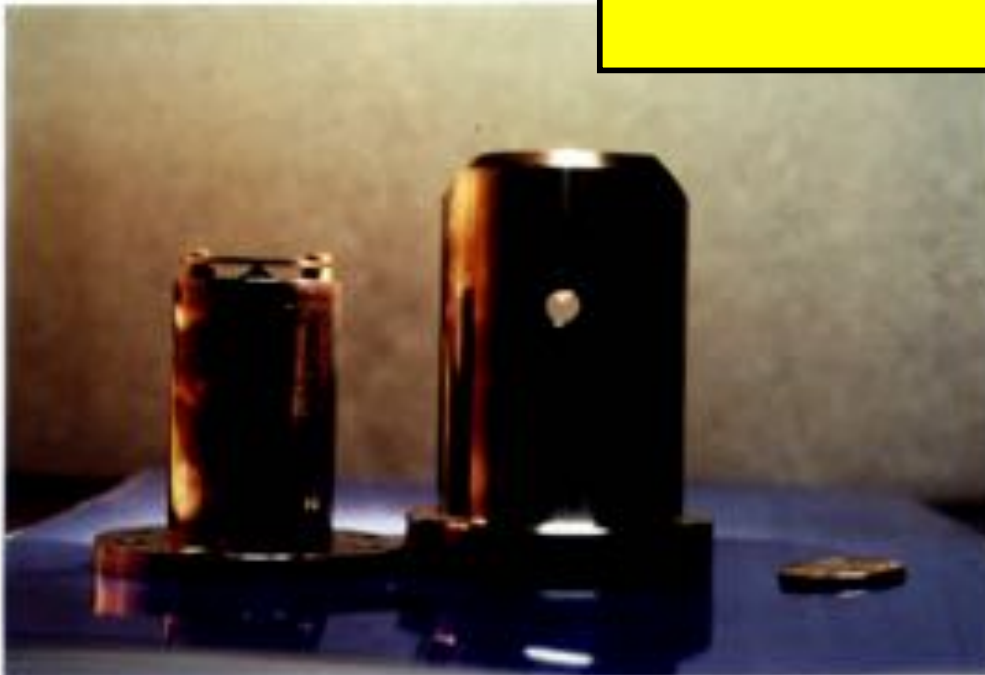
[^]All Fe is listed as FeO.



DIAMOND ANVIL CELL

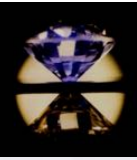


$$\text{Pressure} = \frac{\text{Force}}{\text{Area}}$$



Diamond

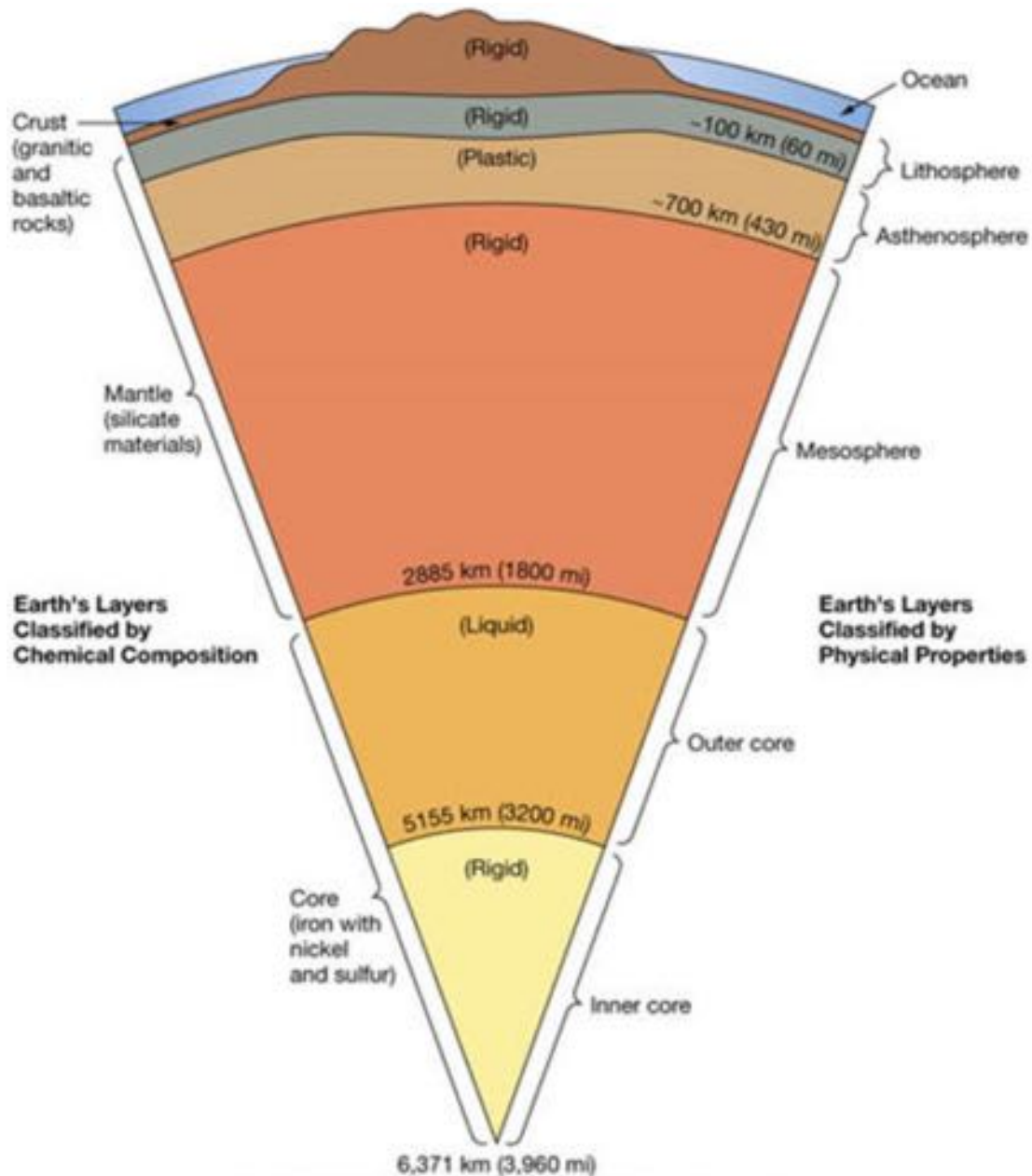
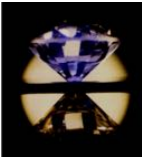
- Strength
- Transparency



~700 African elephants → highest pressures in the Earth



Earth's current P/T conditions



0 GPa, 300 K

24 GPa, ~2000 K

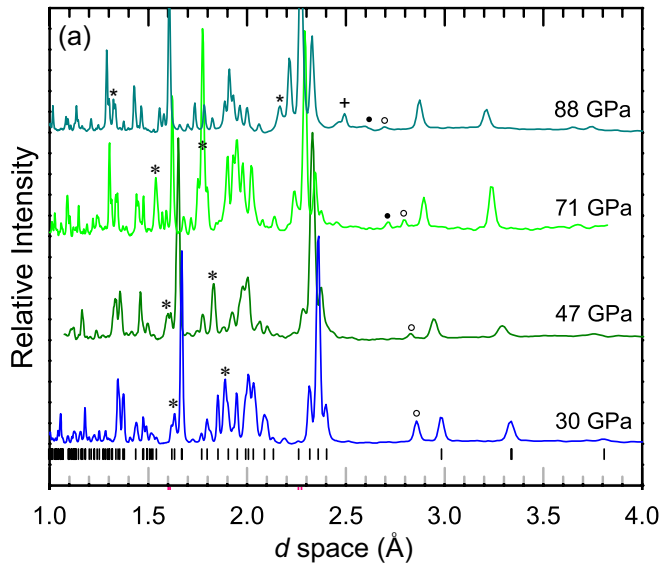
136 GPa, ~4000 K ?

330 GPa, ~6000 K ?

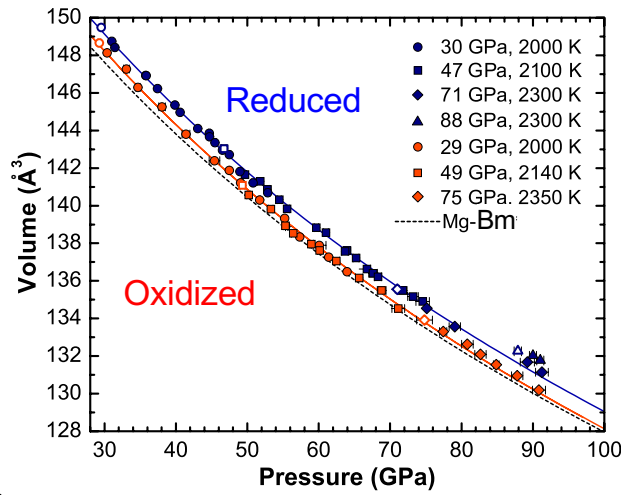
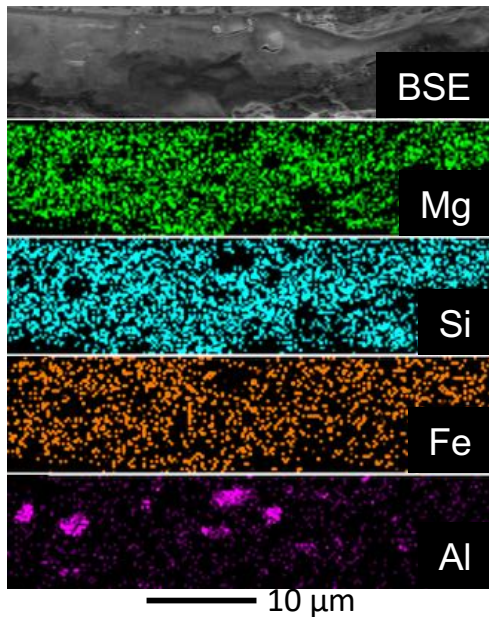
364 GPa, ~6000 K

Effect of Redox on Mantle Mineralogy: J95

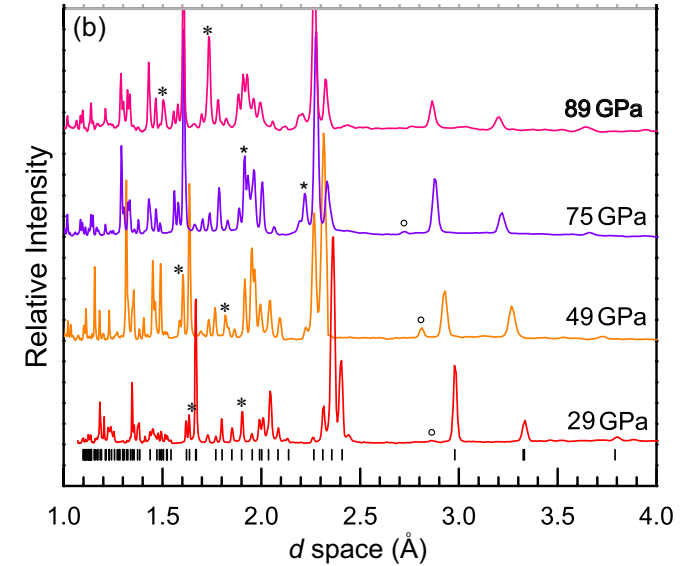
Reduced En Chon LM ($\text{Fe}^{3+}/\text{Fe} \sim 30\%$):



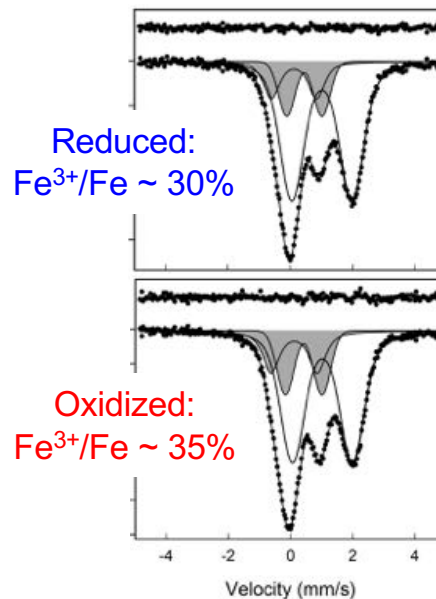
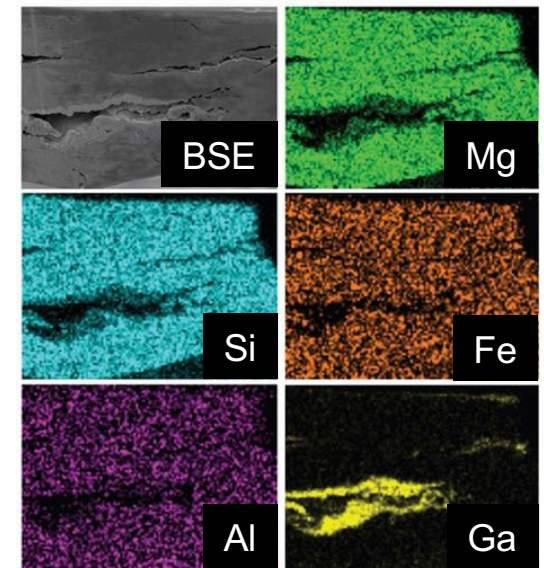
94% Bridgmanite, 5% Silica
1% Alumina



Oxidized En Chon LM ($\text{Fe}^{3+}/\text{Fe} \sim 35\%$):

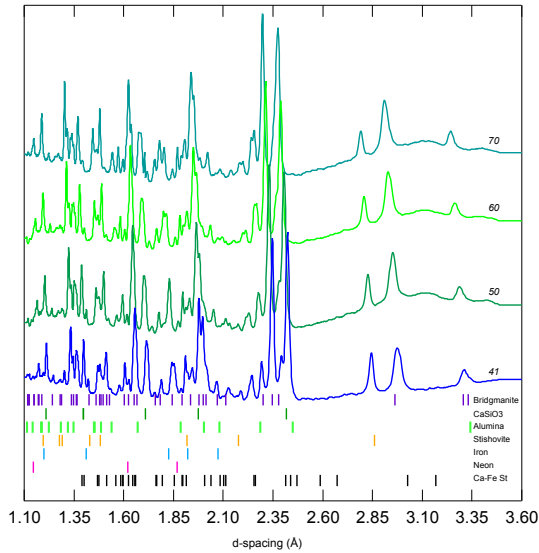


98% Bridgmanite, 2% Silica

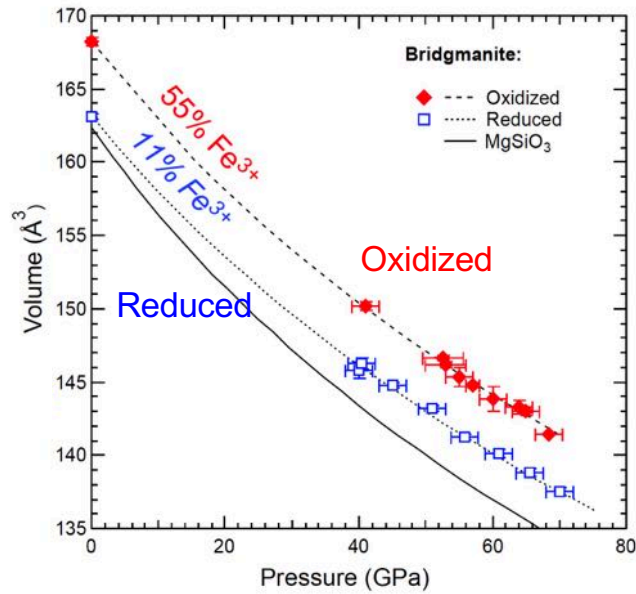
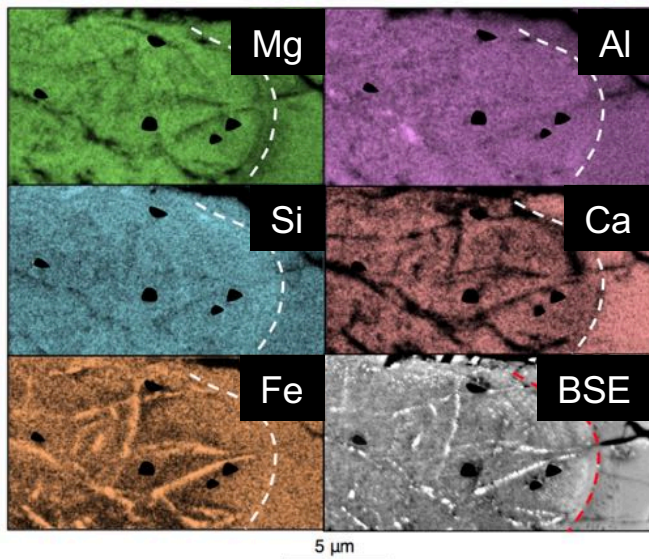


Effect of Redox on Mantle Mineralogy: MIX1G

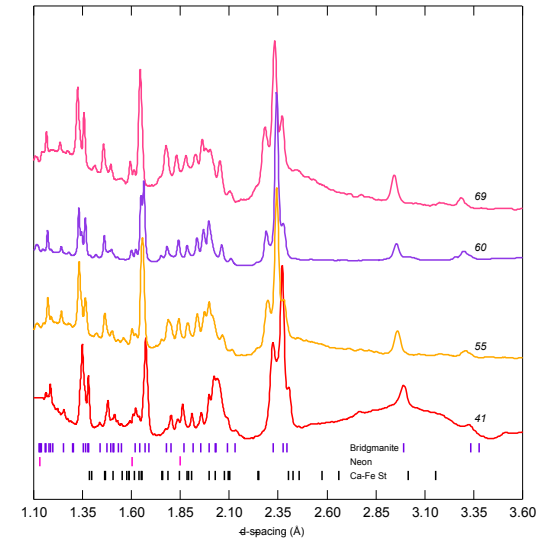
Reduced Pyroxenite ($\text{Fe}^{3+}/\text{Fe} \sim 11\%$):



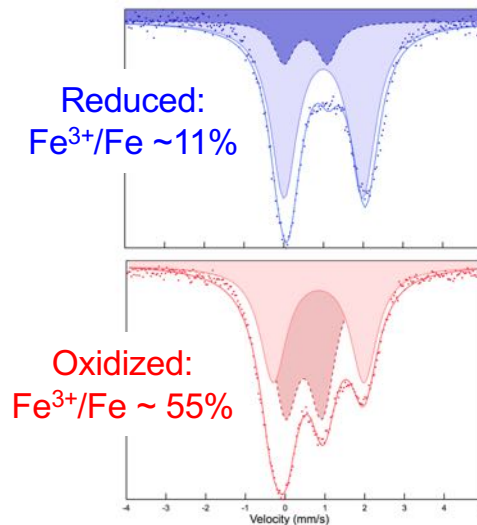
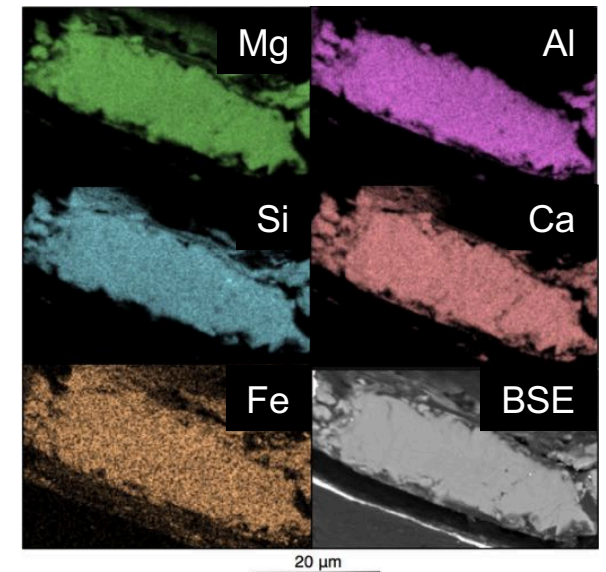
Bridgmanite (51%), Calcium perovskite (18%), CF (14%), Alumina (11%), Stishovite (5%), Iron (1%)



Oxidized Pyroxenite ($\text{Fe}^{3+}/\text{Fe} \sim 55\%$):



Bridgmanite (100%)



Creasy et al., under review

Monte Carlo Modeling Mantle Mineralogy

Assumptions used for the reduced MIX1G samples

6 possible phases: Bm, Capv, Cf, Al, Stv, and Fe

Si bearing phases: Bm, Capv, Stv

Mg bearing phases: Bm, Cf

Al bearing phases: Bm, Cf, Al

Fe⁰ bearing phases: Fe

Fe²⁺ bearing phases: Bm, Cf

Fe³⁺ bearing phases: Bm

Ca bearing phase: Capv

Unknowns

Distribution of Al between Bm, Al and Cf

Distribution of Si between Bm and Stv

Distribution of Mg between Bm and Cf

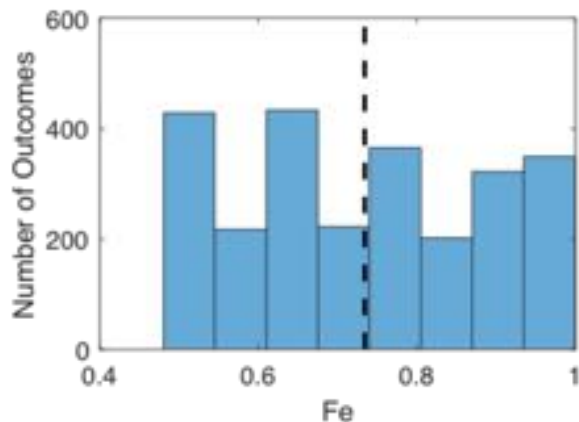
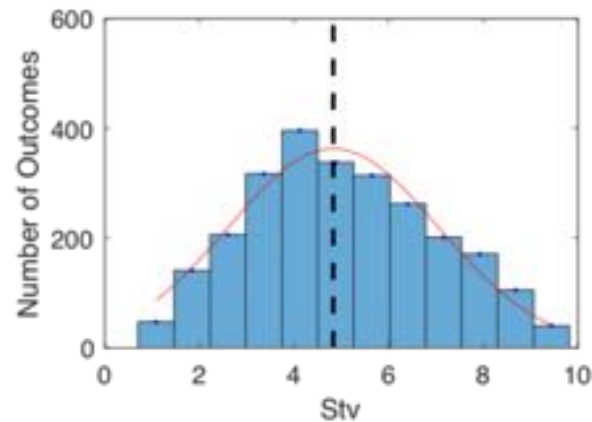
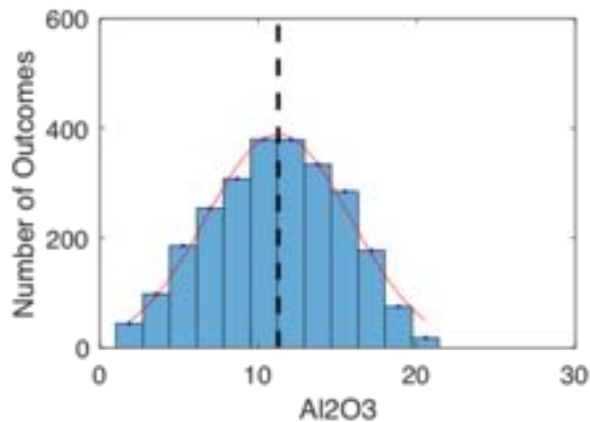
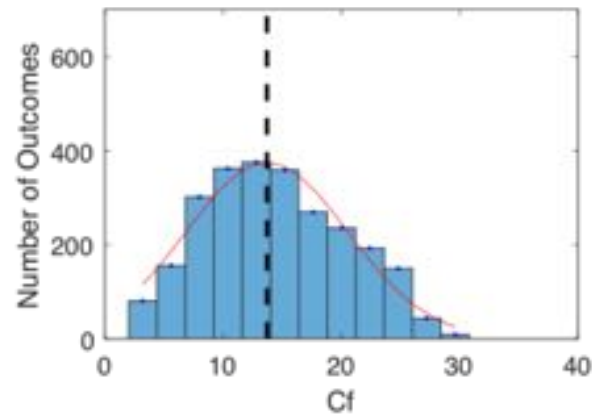
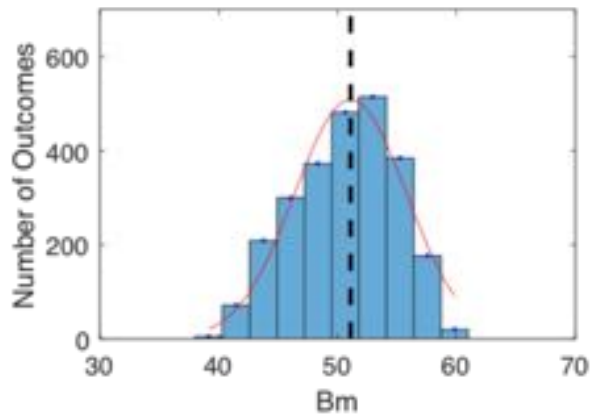
Distribution of Fe²⁺ between Bm and Cf

Amount of Fe²⁺ disproportionation: $3\text{Fe}^{2+} = 2\text{Fe}^{3+} + \text{Fe}^0$

Constraints

- (1) Each mineral is charge balanced
- (2) Self-reduction of iron occurred (we observe the presence of metallic Fe)
- (3) Capv is the only calcium-bearing phase for the MIX_RED samples
- (4) Cf phase is a solid solution between MgAl_2O_4 and FeAl_2O_4
- (5) Use the measured Bm V_0 and existing literature to estimate volume expansion due to cation incorporation
- (6) Assume all Fe³⁺ is incorporated in Bm
- (7) Total Fe³⁺ content cannot exceed 55%
- (8) Metallic iron is present between 0.5 to 1.0% based on EPMA results

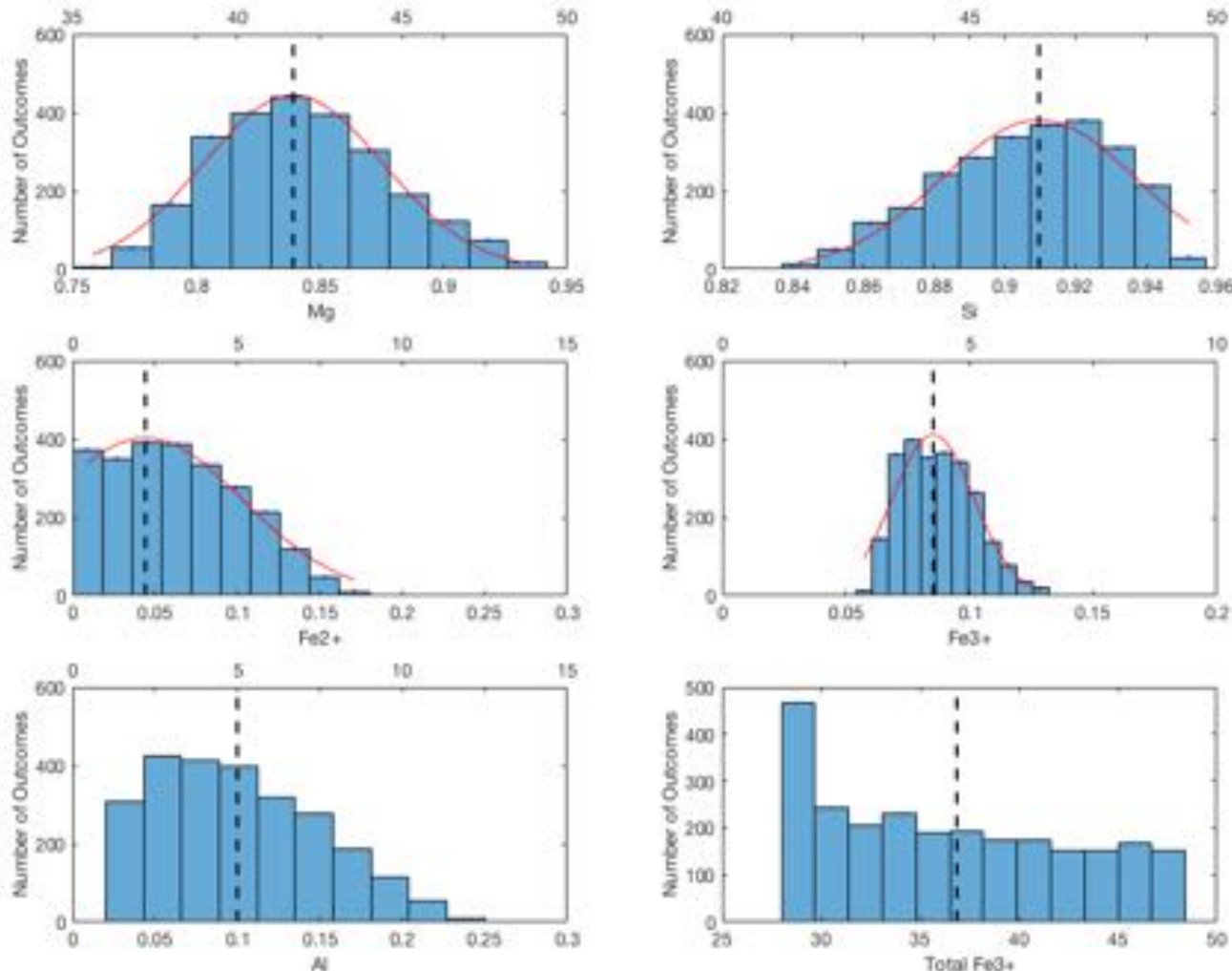
Monte Carlo Modeling Mantle Mineralogy, II



Monte Carlo results for estimating the composition of the synthesized MIX_RED: ~3,000 possible compositions from 3 million iterations

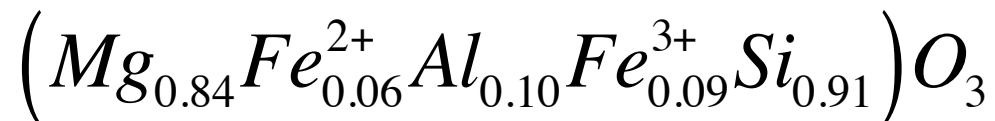
Bridgmanite (51%), Calcium perovskite (18%), CF (14%), Alumina (11%), Stishovite (5%), Iron (1%)

Monte Carlo Modeling Mantle Mineralogy, III



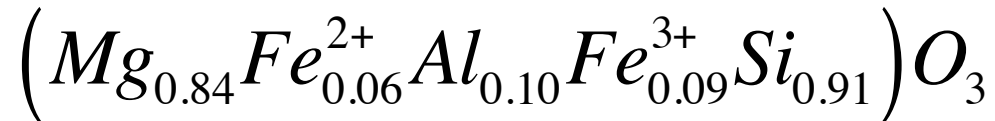
Monte Carlo results
for estimating the
composition of the
synthesized
MIX_RED: ~3,000
possible
compositions from
3 million iterations

MIX_RED Bridgmanite composition:



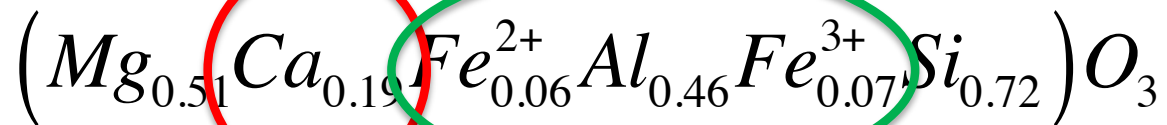
Comparing Bridgmanite Compositions

MIX_RED Bridgmanite composition:



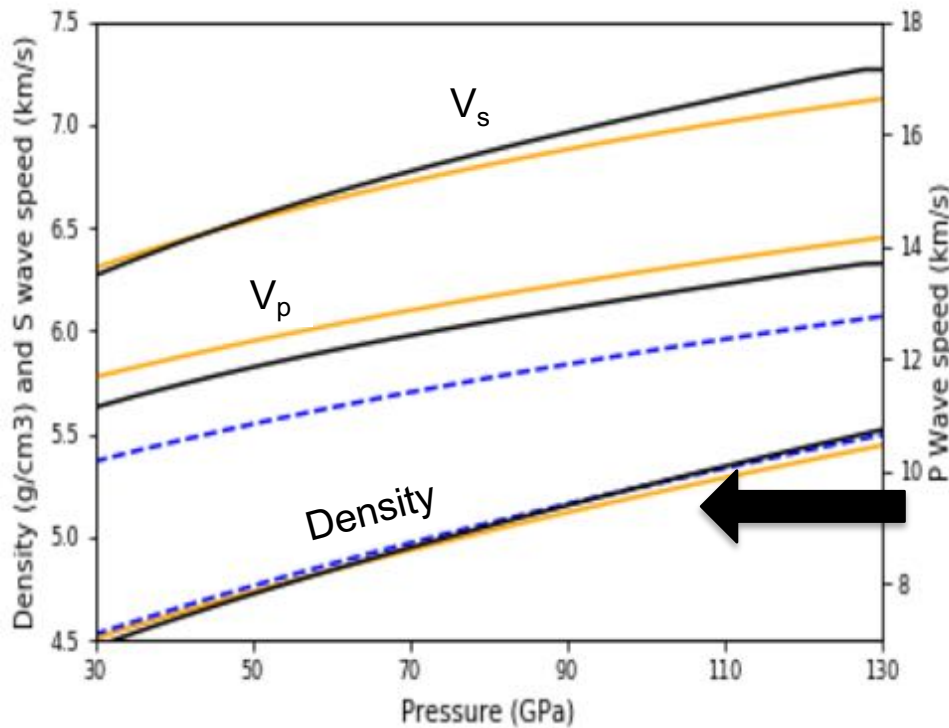
$$V_0 = 163.6 (\pm 0.5) \text{ \AA}^3$$

MIX_OX Bridgmanite composition:



$$V_0 = 168.2 (\pm 0.5) \text{ \AA}^3$$

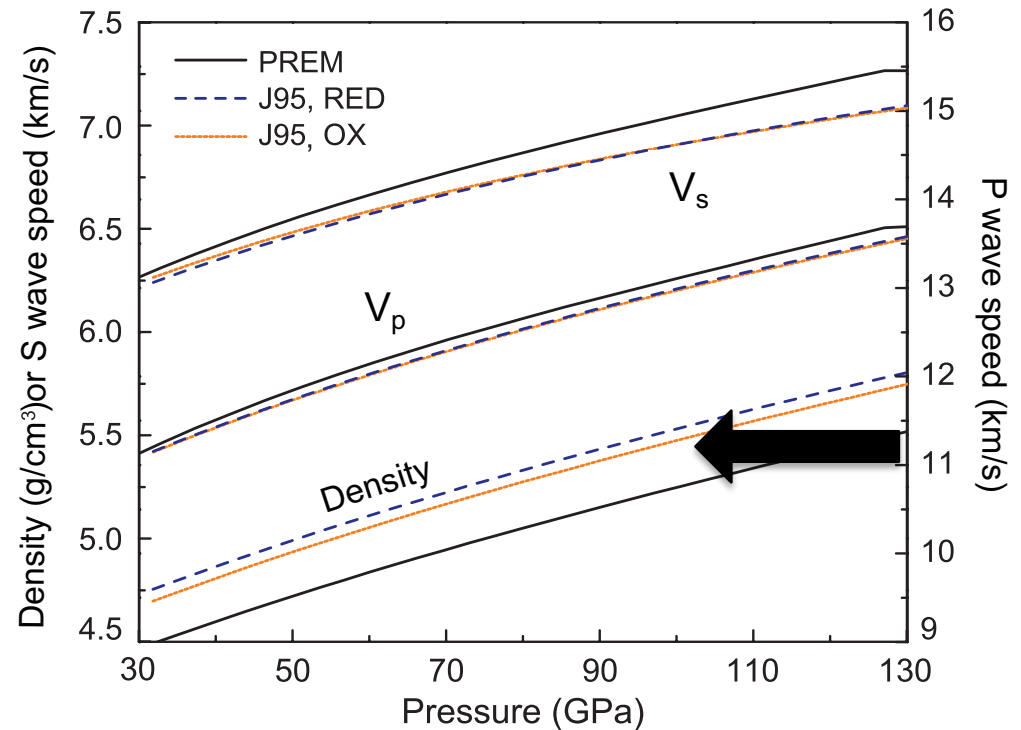
Comparison with PREM



MIX1G

~1% density difference
due to redox state!

Reduced: $\text{Fe}^{3+}/\Sigma\text{Fe} \sim 11\%$
Oxidized: $\text{Fe}^{3+}/\Sigma\text{Fe} \sim 55\%$

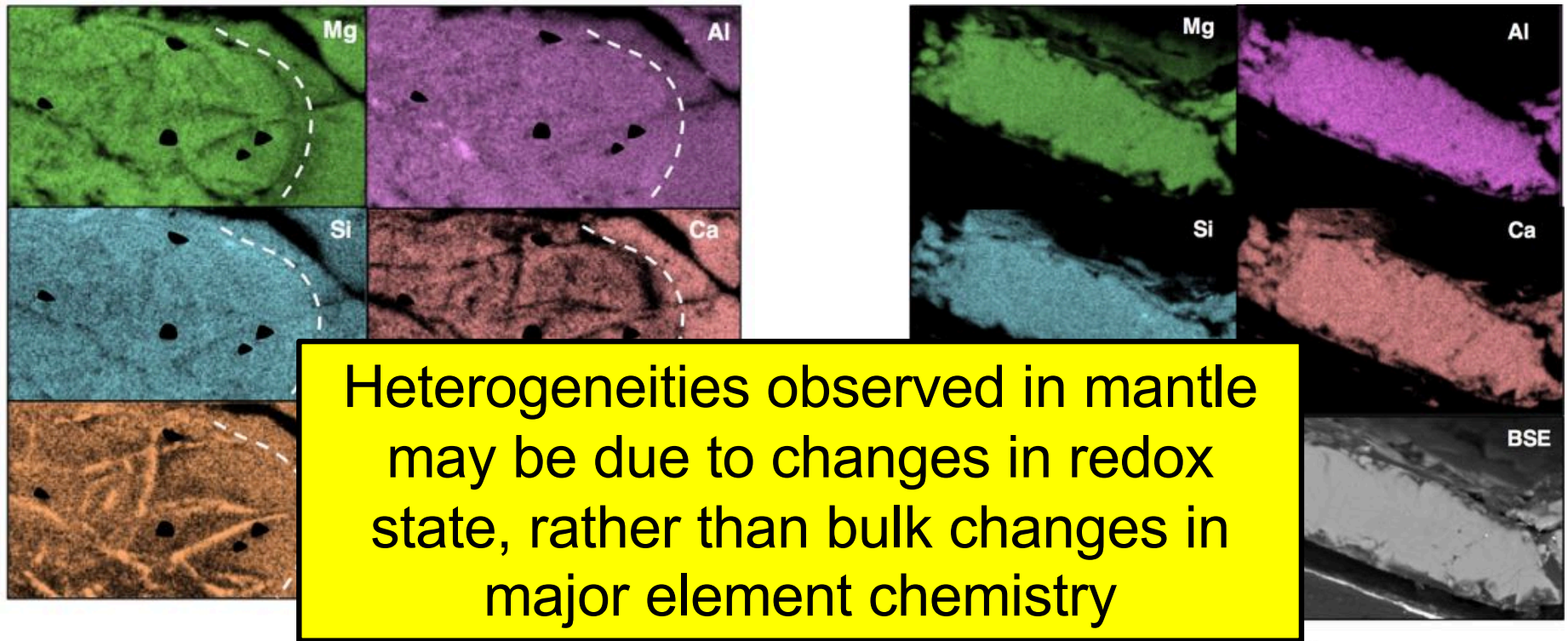


J95

~1.5% density difference
due to redox state!

Reduced: $\text{Fe}^{3+}/\Sigma\text{Fe} \sim 30\%$
Oxidized: $\text{Fe}^{3+}/\Sigma\text{Fe} \sim 35\%$

Effect of Redox on Mantle Mineralogy



Reduced samples:

1. More complex mineralogy
2. Greater assemblage density

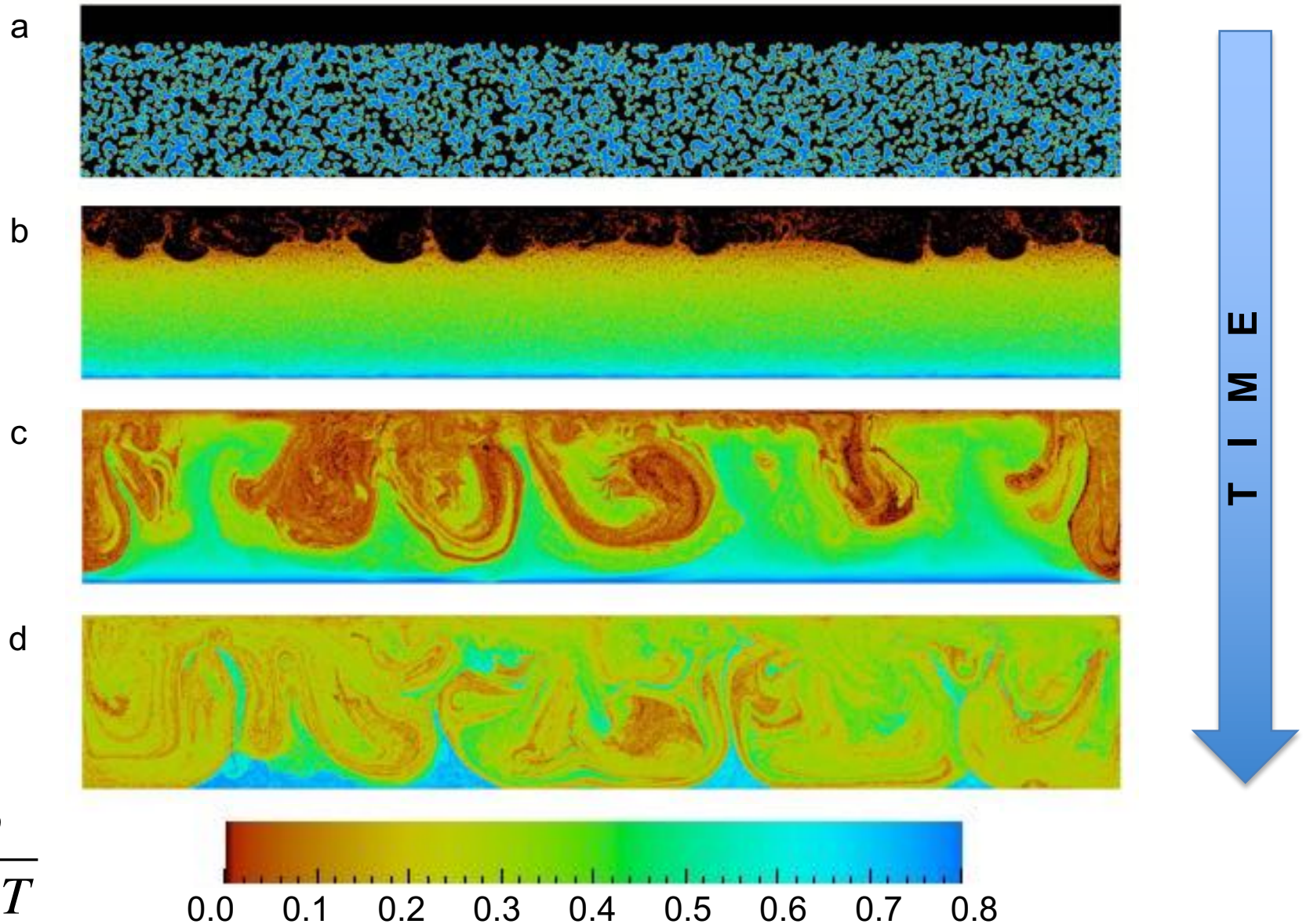
Oxidized samples:

1. Mostly bridgmanite
2. Lower assemblage density
3. Solid solution between Mg- and Ca-silicate perovskites

Geodynamic Modeling

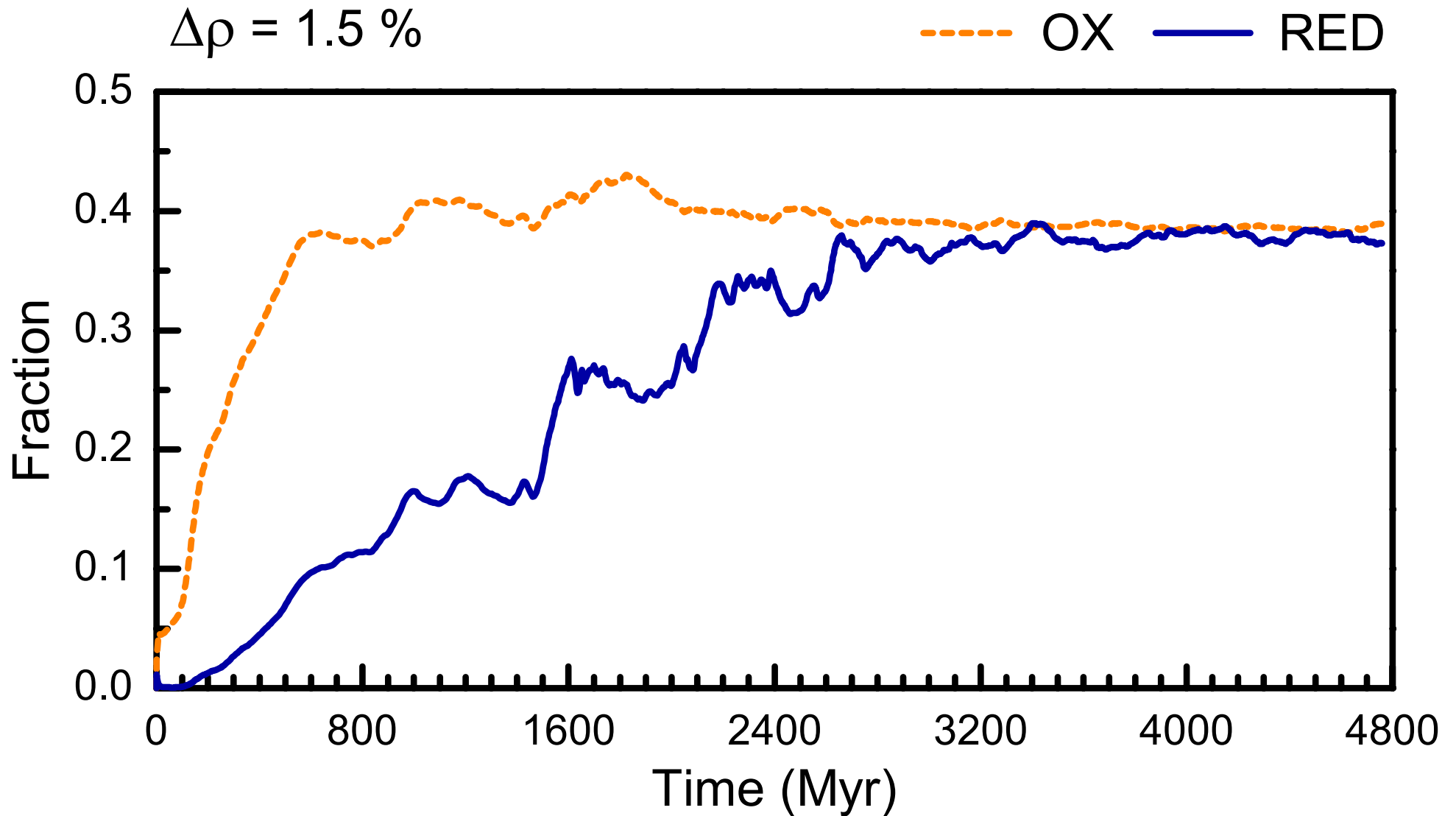


Geodynamic Modeling

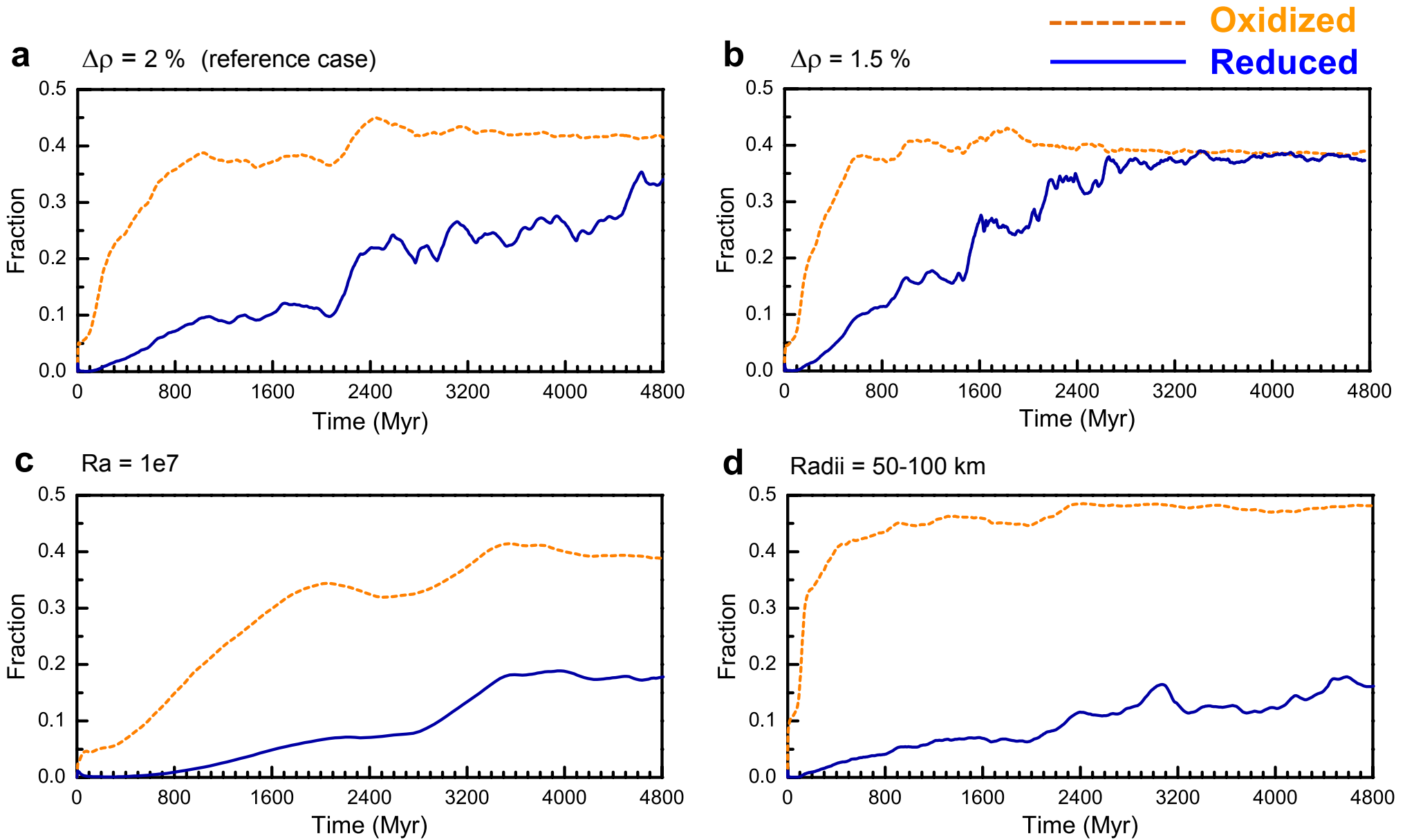


$$B = \frac{\Delta\rho}{\rho\alpha\Delta T}$$

Evolution of upper mantle



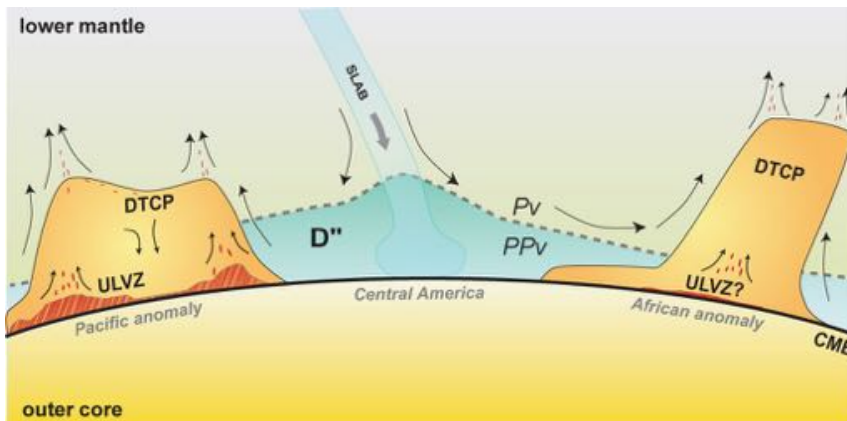
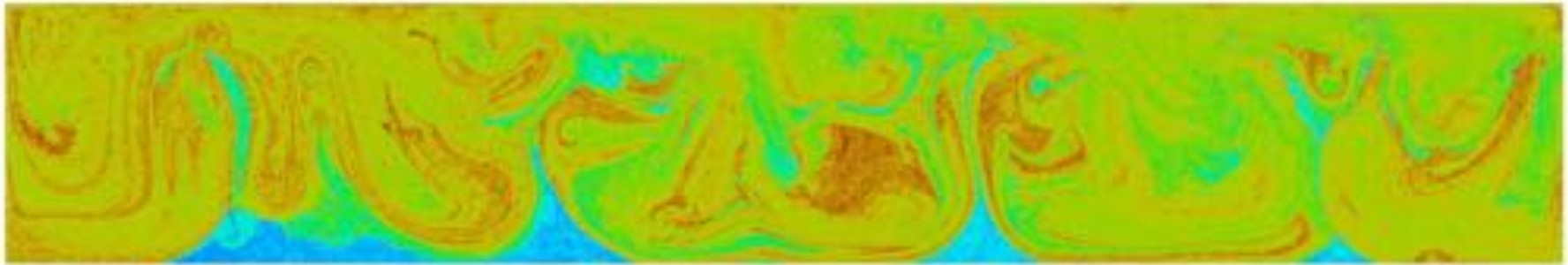
Evolution of upper mantle



What could this all mean?



Lots of O₂ in the atmosphere!

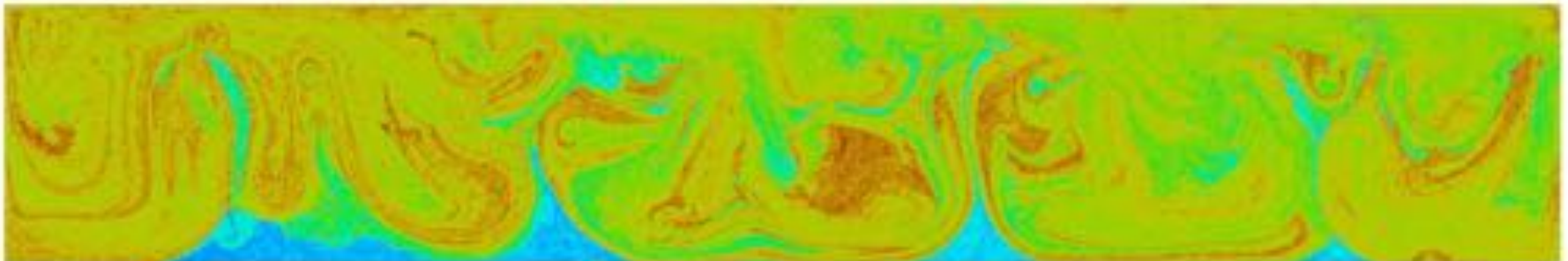


[Garnero et al. (2006)]

Reduced, dense,
primordial piles?

Summary

- Crystal chemistry in bridgmanite varies under different redox conditions influencing mantle mineralogy, density, seismic velocities and convection.
- More oxidized samples (high $\text{Fe}^{3+}/\Sigma\text{Fe}$), yield simpler assemblages and are less dense than their more reduced counterparts by ~1-2%.
- While there are several possible chemical models of the lower mantle, bridgmanite is nevertheless considered the dominant phase in the lower mantle; thus, its crystal chemistry has a leading role in the thermochemical evolution of the mantle (*and exosphere!*).



What comes next?

- How does Ca incorporation into bridgmanite affect partitioning of other large-ion lithophile elements (e.g., K, U, Th)?
- How are melting relations affected by redox state?
- What are the observables that can further test this hypothesis?
- What might this mean for planets beyond Earth, both within and outside of the Solar System?
- ...

